

## AVAILABILITY OF SOLAR RADIATION AND ITS ESTIMATION

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**Keywords:** Albedo, Clearness of sky, Ecliptic, Electromagnetic spectrum, Extraterrestrial radiation, Incident radiation, Insolation

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### Summary

Solar energy is spread out and is available in all places around the world. Its intensity varies significantly according to the place on the earth's surface and for the same place according to the time of the day, the time of the year and the existing meteorological conditions. It is very spread out, in dilute concentration. As a very important renewable energy source solar radiation parameters and solar-earth angles are of importance for the estimation of the energy reaching the earth's surface.

## **1. Introduction**

Solar energy is emitted in huge amounts from the sun's surface but only a fraction of that reaching the earth's surface can be converted into useful forms of energy for utilization either as thermal, electrical or mechanical. For desalination purposes solar energy can be used either directly to distill water in solar stills or can be utilized indirectly to drive conventional desalination plants. In indirect use of solar energy solar radiation has to be collected and converted into either heat or electricity. The transformed solar radiation can then be applied to drive one of the conventional desalination systems. Thus in an indirect utilization of solar energy two separate plants are involved: a solar thermal or electricity plant and the desalination plant. Collectors or concentrators are used to convert the radiation into thermal energy in the form of hot water or steam. The steam generated is of low or high temperature, depending from the type of the solar collector field. Low and medium temperature steam is used as process heat in industry, including desalination plants, or as space heating. Electricity is produced either by high temperature steam in solar power systems or by solar cells in photovoltaic fields.

The irregularity of solar radiation intensity affects the normal and smooth operation of the desalination plants, which operate in a discontinuous and unsteady state basis if storage is not provided. Storage devices are necessary for continuous operation, during night time or during cloudy days, but are also capital intensive enterprises adding to the already high cost of solar energy conversion systems. Although there are these disadvantages, thermal solar energy systems are suitable for providing energy to medium or small size desalination plants in sites where there is abundant and intensive solar radiation. Normally all solar powered plants offer energy, at a reasonable cost, suitable to use in desalination facilities.

Solar energy may be available in some form everywhere but it is not available at all times anywhere on the earth. To determine whether and how solar energy is available and, or, usable on an economic basis precise meteorological data must be known, and all data concerning the incident solar radiation of the region. To operate all solar plants and desalination solar systems, the spatial and the temporal, i.e. the hourly, daily and seasonal, variability data of solar radiation is of importance. These data are also indispensable for the proper design of a solar driven system, and for the proper selection of devices and methods to be used. Their most important utilization is in the evaluation of the performance of both plants. A detailed understanding of the spatial, temporal and the spectral characteristics of the solar radiation is attained through detailed and exact measurements, detailed analysis of the data selected, and computer modeling of the variations in solar radiation. It is thus very important for an engineer dealing with solar desalination to know all about the sun and its global radiation emitted around the year.

## **2. The Sun and the Radiation Emitted**

### **2.1. The Physics of the Sun**

The sun is a sphere consisting of intensively hot ionized gaseous matter, called plasma. In fact the sun is a large nuclear reactor where thermo-nuclear fusion reactions take

place continuously generating huge amounts of energy. The energy radiated into the surrounding space is characterized by the sun's structural characteristics. These principal characteristics (Sayigh 1984) are:

Linear diameter	$1.392 \times 10^6$ km
Sphere mass	$(1.991 \pm 0.002) \times 10^{30}$ kg
Average mass density	$(1.410 \pm 0.002) \times 10^3$ kg m <sup>-3</sup>
Effective blackbody surface temperature	$5762 \pm 50$ K
Interior temperature	$8 \times 10^6$ to $40 \times 10^6$ K
Average distance from earth	$1.495 \times 10^6 \pm 1.7\%$ km
Core density	$100 \times 10^3$ kg m <sup>-3</sup>
Surface density	$\sim 10^{-5}$ kg m <sup>-3</sup>
Solar energy at the Earth atmosphere	$178 \times 10^{14}$ kW
Time for the ray to reach Earth	8 min

The linear diameter corresponds not directly to the real diameter, as the sun's atmosphere extends far beyond the photosphere, with decreasing density. The thermonuclear processes that take place in the mass of the sun lead to the production of radiation particles called photons. In the vicinity of the earth the atmosphere of the sun has  $100 \times 10^6$  to  $400 \times 10^6$  protons per m<sup>3</sup> and about  $0.7 \times 10^6$  hydrogen atoms per m<sup>3</sup>. Photons are energy units having zero charge and no mass. The energy of a photon  $E$  is proportional to its frequency and the Planck's constant,  $h$  (with numerical value =  $6.6 \times 10^{-34}$  J).

$$E = h \times \nu = \nu \times 6.6256 \times 10^{-34} \text{ J} \quad (1)$$

The photon energy increases with frequency increase, i.e. increases with decreasing of the corresponding wavelength. When the thermonuclear processes are in progress in the sun, molecules, atoms and electrons are raised to excited states and then momentarily return to low energy states releasing thermal energy known as electromagnetic radiation. It is estimated that about 90% of the energy emitted is generated in the region of 0 to 0.23 of its radius, a region where about 40% of the sun's mass is concentrated. The energy released from the core is transferred upwards, in the form of X-rays and  $\gamma$ -rays, to the upper zone of the sun near the surface. This zone consist of an opaque strongly ionized gas layer, called "photosphere", which absorbs and emits a continuous spectrum of radiation. It is the source of light and heat radiated to the earth.

## 2.2. The Electromagnetic Spectrum

The colossal fusion that takes place in the sun' mass releases tremendous quantities of electromagnetic radiation, as result of its mass fusion. Calculations derived from Einstein's mass-energy conversion equation,  $E = mc^2$ , have shown that the sun loses  $4 \times 10^9$  kg s<sup>-1</sup>, or  $0.5 \times 10^{-21}$  of its total mass per second. This amount of mass is released as energy and it is distributed into space in the form of electromagnetic radiation. Its spectrum varies from very short wave length cosmic radiation to very high frequency long wavelength radiation, i.e. from fractions of Ångstrom ( $10^{-10}$  m) to hundreds of

meters. The spectrum is divided into wavelength bands which travel into space with the speed of light,  $c$ , and is connected to the wavelength  $\lambda$  and the frequency  $\nu$  by the following simple equation:

$$c = \lambda \times \nu = 2.99776 \times 10^8 \text{ m s}^{-1} \quad (2)$$

The approximate wavelength bands are given by Robinson (1966) as follow:

10 Å X and $\gamma$ rays	7200 Å-1.5 $\mu\text{m}$ near infrared
10-2000 Å	Far ultraviolet 1.5-5.6 $\mu\text{m}$ middle infrared
2000-3150 Å	Middle ultraviolet 5.6-1000 $\mu\text{m}$ far infrared
3150-3800 Å	Near ultraviolet > 1000 $\mu\text{m}$ micro
3800-7200 Å	Visible region radio-waves

The spectrum between 2000 and 30 000 Å carries 98% of the total emitted energy and it is called the "quiet sun" region. The temperatures in this region are  $8 \times 10^6$  to  $40 \times 10^6$  K. Wave lengths greater than 30 000 Å, belong to the infrared region and are absorbed by water vapor and carbon dioxide in the atmosphere. In the ultraviolet region wavelengths of approximately 2860 Å reach the sea level but shorter wavelengths are absorbed by the ozone layer in the upper regions of the atmosphere. The solar radiation that reaches the earth's surface consists of wave lengths of the ultraviolet to near infrared range, i.e. from 0.3 to 2.5  $\mu\text{m}$ . For most solar applications the radiation in the visible range (0.38 to 0.78  $\mu\text{m}$ ) and the near infrared (0.78 to about 2  $\mu\text{m}$ ) is of most importance.

The sun is considered as a black body, thus the total radiation emitted per unit area from its outer surface, the photosphere, is calculated from the Stefan-Bolzman equation:

$$G_s = \sigma \times T^4 = 5.67 \times 10^{-8} (5762)^4 = 62.50 \times 10^3 \text{ kWm}^{-2} \quad (3)$$

Wavelength range $\lambda$ [ $\mu\text{m}$ ],	0-0.38	0.38-0.78	0.78-40
Approximate energy $\text{W m}^{-2}$	95	640	605
Approximate % of total energy	7.0%	47.3%	44.7%

Table 1. The solar energy spectrum distribution (Mangal 1990).

Taking in consideration the total surface of the sun, for  $D_s = 1.392 \times 10^9$  m, the total power emitted by the sun is given as:

$$\pi \times (D_s)^2 \times G_s = 3.14159 \times (1.392 \times 10^9)^2 \times 62.50 \times 10^3 = 3.8046 \times 10^{23} \text{ kW} \quad (4)$$

Out of this amount the earth intercepts only  $1.78 \times 10^{14}$  kW. The amount of electromagnetic radiation emitted from the sun in the main wavelength regions is given in Table 1.

### 3. The Extraterrestrial Radiation

#### 3.1. What is the Solar Constant

The incident spectral solar radiation outside the earth's atmosphere is called "extraterrestrial",  $G_o$ , or, air-mass zero (AMO) solar radiation. Its instantaneous power the "irradiance" or, solar flux density is measured in W per square meter ( $\text{W m}^{-2}$ ) and is considered as a constant, although there exist some fluctuations due to solar activities. At the top of the atmosphere and at a mean earth-sun distance  $r_o$ , the intensity of this radiation is termed as the "Solar Constant",  $G_{sc}$ , which is defined as: "the perpendicular radiation that receives a surface of one square meter, at the earth's mean distance from the sun per unit of time". Many scientists give in their studies values of the Solar Constant, calculated from terrestrial measurements. Today most measurements are performed in extraterrestrial space by satellites.

NASA (1971) performed extraterrestrial measurements by spacecraft and as a result the exact value of the solar constant was calculated as  $G_{sc} = 1353 \pm 0.021 \text{ W m}^{-2}$ , or  $4.871 \pm 75.5 \text{ kJ m}^{-2} \text{ h}^{-1}$ . Later measurements by Iqbal (1983) give a value for the solar constant of  $1367 \text{ W m}^{-2}$  or  $4.921 \text{ MJ m}^{-2} \text{ h}^{-1}$ . Today solar constant values are derived from satellite and spacecraft measurements with differences less than 1.0 per cent. The World Radiation Center (WRC) adopted the Iqbal value accepting an uncertainty of 1.0 per cent. It is the value of solar constant accepted by the most scientists in their scientific work.

#### 3.2. Variations of the Extraterrestrial Radiation

The earth has an elliptical orbit that results in approximately  $\pm 3.3\%$  variation in the amount of solar radiation at the top of the atmosphere throughout the year. For a plane,  $p$ , outside the atmosphere, at a distance,  $r$ , from the earth, the extraterrestrial radiation is given as:  $G_o = G_{sc}(r_o/r)$ . The extraterrestrial intensities for the 21<sup>st</sup> of each month of the year are presented in Table 2. They can be also calculated from a set of equations as a function of the sun's angles, which are defined in (Section: Flat-plate collectors). The extraterrestrial solar radiation and the solar constant are the initial input data for the solar radiation calculations that receive solar collectors. They are used to formulate equations for the scattering and the absorption processes taking place in the atmosphere and also to calculate the "Clearness Index". The extraterrestrial radiation  $G_o$  is calculated as a function of the solar constant and a random day of the year from Eqs (14) and (15) given below.

#### 3.3. The Black-body Emittance and Absorptance

The sun acts as an effective radiation blackbody, independent of the incident directions. A perfect absorber is also a perfect emitter of radiation, thus it emits the maximum possible radiation. It has to absorb and emit all spectral radiation maintaining a constant temperature. The blackbody is an imaginary concept, because no real material is a perfect absorber or emitter. Nevertheless some materials, mainly black in color, approach the blackbody characteristics having an absorptance or emittance of about 99 per cent.

Month	$Q_o$ ( $W m^{-1}$ )	Month	$Q_o$ ( $W m^{-1}$ )	Month	$Q_o$ ( $W m^{-1}$ )
January	1422	May	1344	September	1367
February	1401	June	1333	October	1390
March	1387	July	1334	November	1411
April	1364	August	1345	December	1423

Table 2. Extraterrestrial solar radiation intensities normal to the sun on the 21<sup>st</sup> of each month (Sayigh 1986)

The blackbody energy distribution curves shown in Figure 1, are presented by three curves for temperatures of 5000, 6000 and 7000 K. The total radiation in the shaded curve corresponds to the so called "smoothed" solar curve which lies in the visible wavelength region. The spectrum of this curve is cut off on the ultraviolet side, below the 6000 K curve, in the infrared region, its maximum being at 4700 Å.

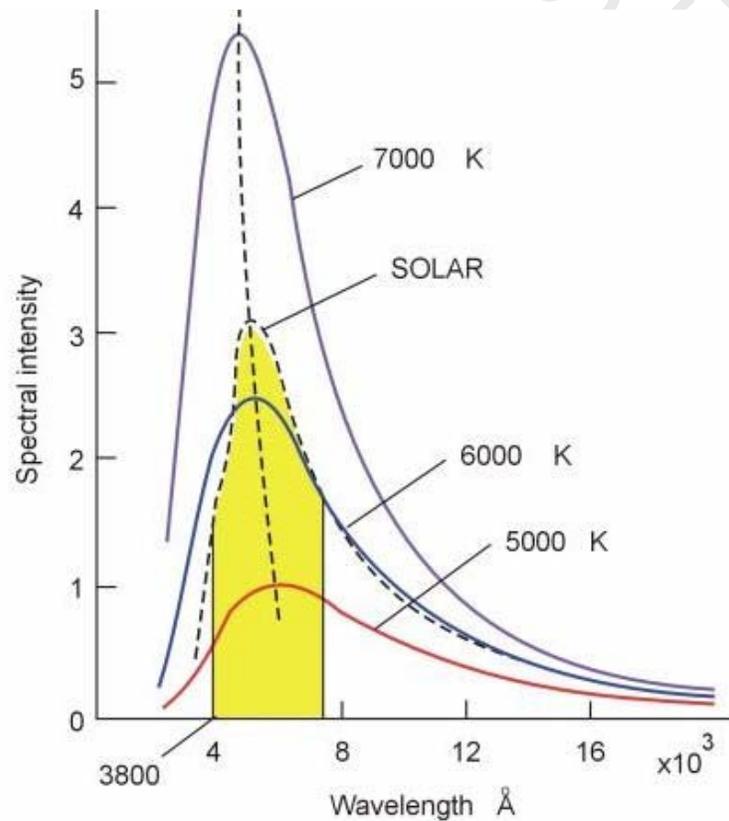


Figure 1. Blackbody energy distribution curves. The sun's spectral curve is presented by the shaded curve and is in the range of visible wavelengths.

### 3.4. Blackbody Absorbers - Application to Desalination

Solar radiation is absorbed in various ratios by physical materials, depending on their special radiation or emittance properties, the perfect absorber being always the

blackbody. By using black material in flat-plate collectors and solar stills, the maximum absorbance of the incident radiation can be reached. The energy  $E_b$  emitted by a perfect blackbody is calculated for all wave lengths  $\lambda$  and for a temperature  $T_b$  K from the Stefan-Boltzmann equation derived from the integration of Planck's law:

$$E_b = E_b \lambda \, d\lambda = \sigma T_b^4 = 5.6697 \times 10^{-8} T_b^4 \, \text{W m}^{-2} \quad (5)$$

For solar distillation and flat-plate collectors efficient sky radiation is an important factor. Sky radiation refers to the radiation exchange between sky and a surface,  $p$ , under observation,  $A_p \, \text{m}^2$ , of temperature  $T_p$  K. The sky is considered as a blackbody of temperature,  $T_s$ . Considering the blackbody perfect absorbance,  $\alpha$ , or emittance,  $\varepsilon$ , as 1.0, all natural materials may have an absorbance or emittance,  $\alpha$ , or,  $\varepsilon < 1.0$ . According to the above Stefan-Boltzmann equation for an emittance,  $\varepsilon$ , of a collector or solar still of absorbance surface  $A \, \text{m}^2$ , the net mutual radiation is given as:

$$Q = \varepsilon \times A_p \times \sigma \times (T_b^4 - T_s^4) = \varepsilon \times A_p \times 5.6697 \times 10^{-8} (T_p^4 - T_s^4) \, \text{J s}^{-1} \quad (6)$$

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